

# Exploring the Key Features of Tropical Cyclogenesis using Satellite Data

Zhuo Wang

Department of Atmospheric Sciences  
University of Illinois at Urbana-Champaign

# Motivating Questions

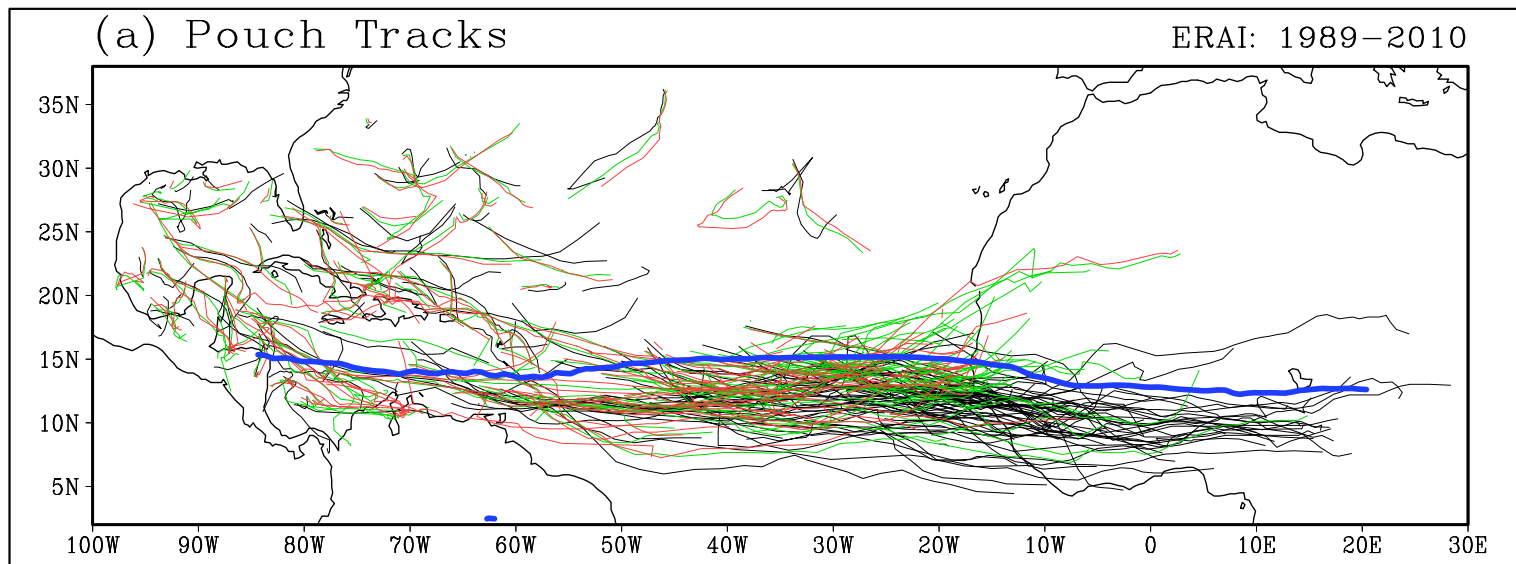
Q1: What is the key feature(s) of convection leading up to genesis?

Q2: What thermodynamic conditions promote the development of such a key feature?

Q3: What are the relative roles of different types of precipitation in TC genesis?

# Data

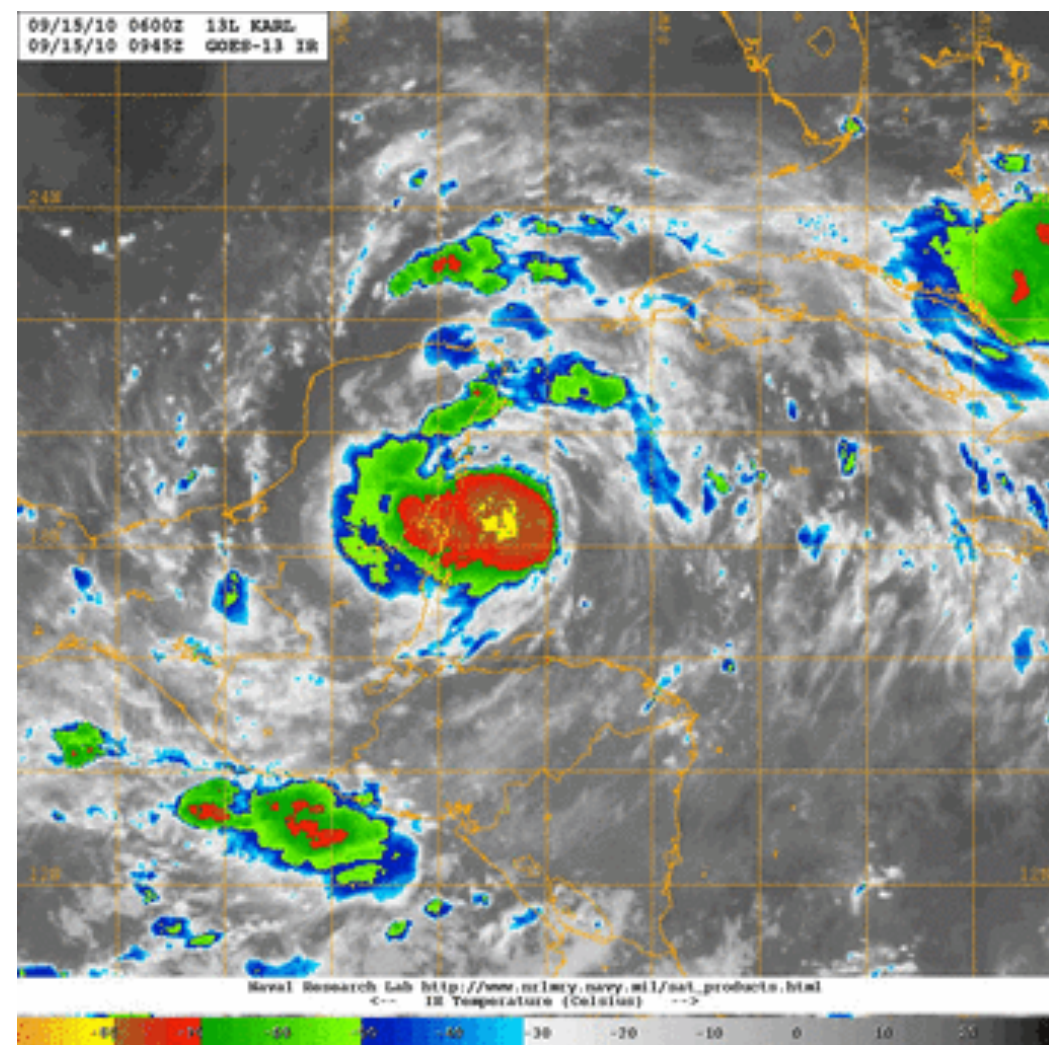
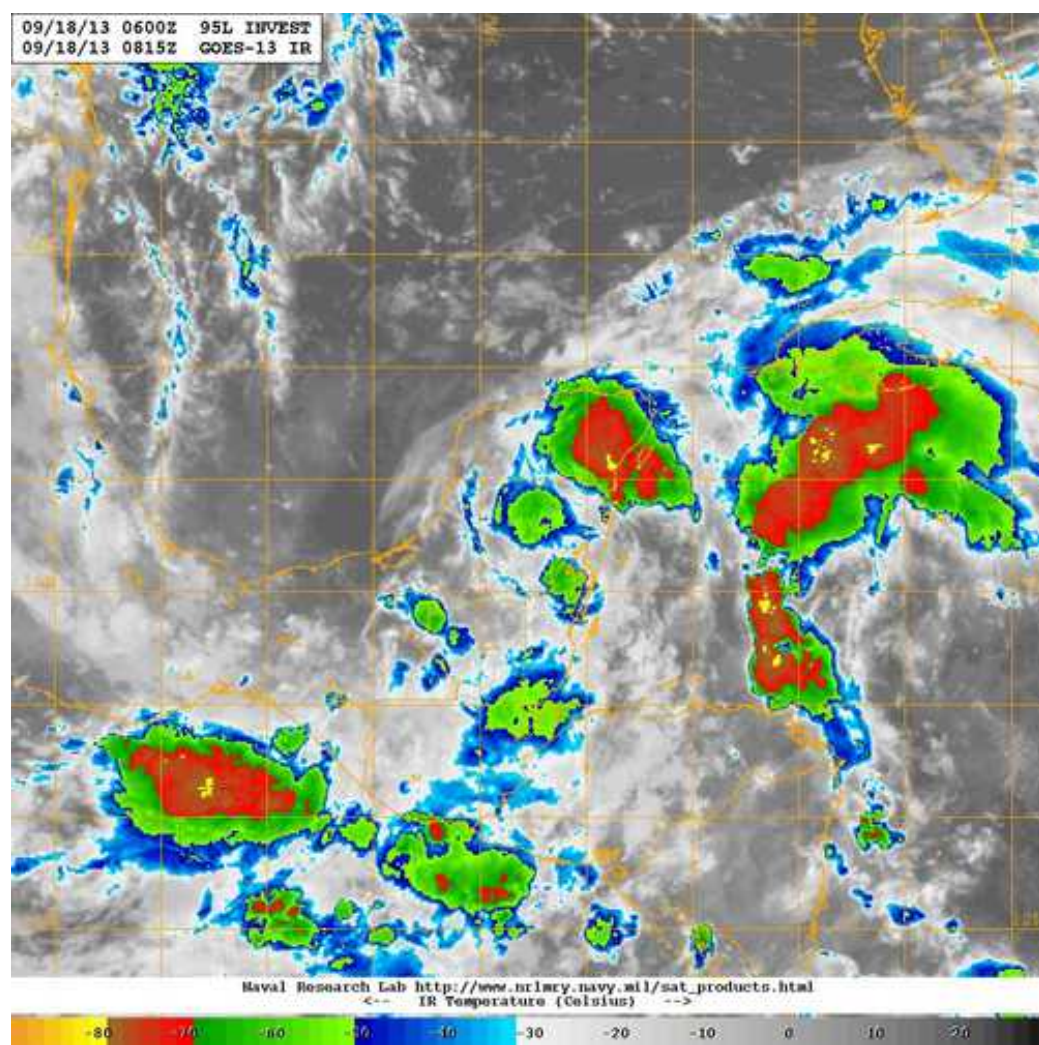
- GribSAT-B1 IR (Knapp et al. 2011)
  - 3-hour data; resolution:  $\sim 7$  km
- Rain rate and column water vapor (CWV) from SSMI/SSMIS
- Rain rate and reflectivity from TRMM PR 2A25
- Best track data to define genesis: the formation of a TD in most cases
- A pouch track dataset for 164 named storms over the Atlantic from 1989-2010 (Wang and Hanks 2014)
  - Like the best track data but for precursors
- 6 hourly data from ERA-Interim Reanalysis



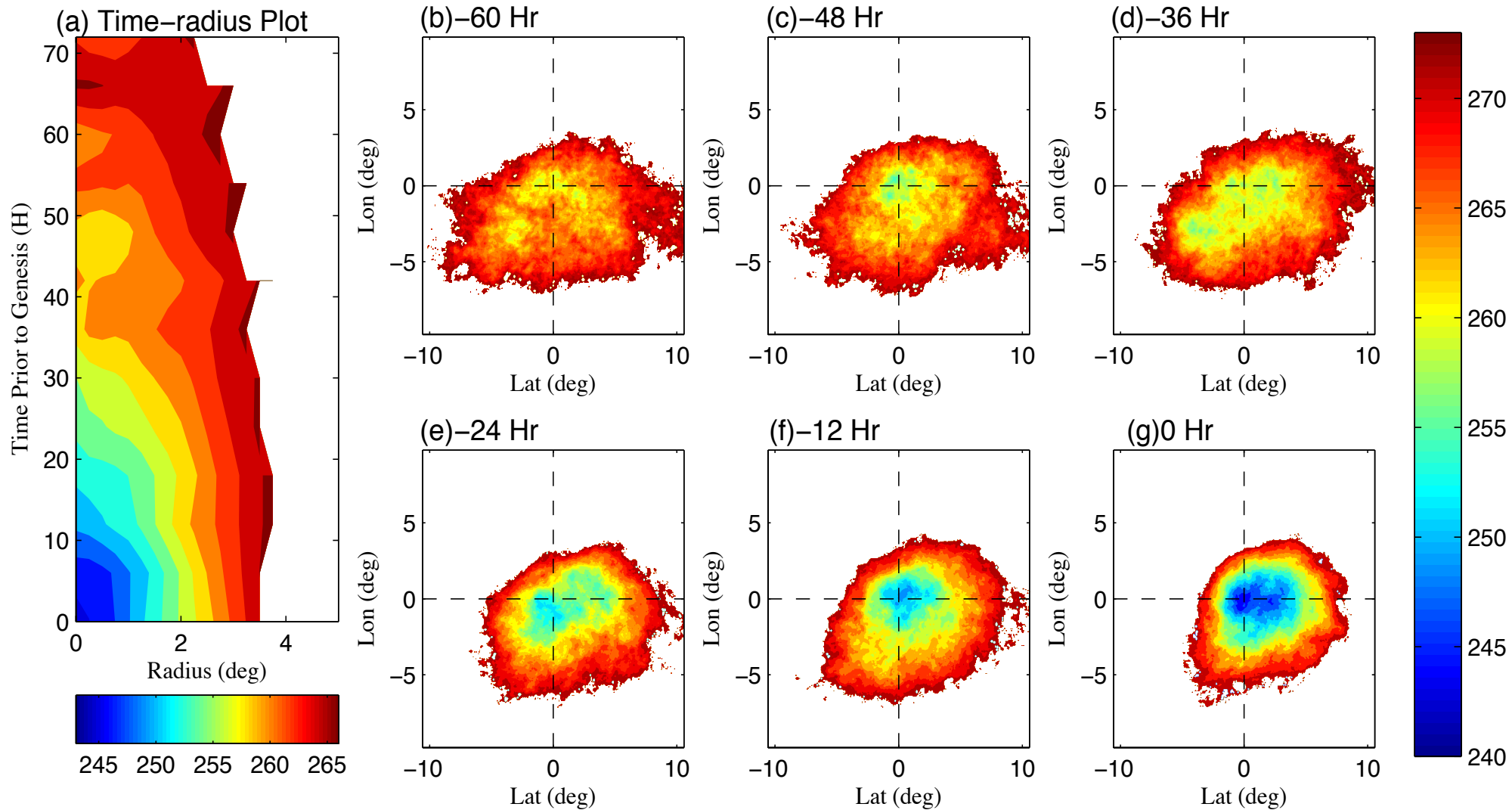
Q1: What is the key feature of convection leading up to genesis?

--- GRIBSAT IR data: high spatial and temporal resolution with continuous observation

# IR imagery: which one is a tropical cyclone?



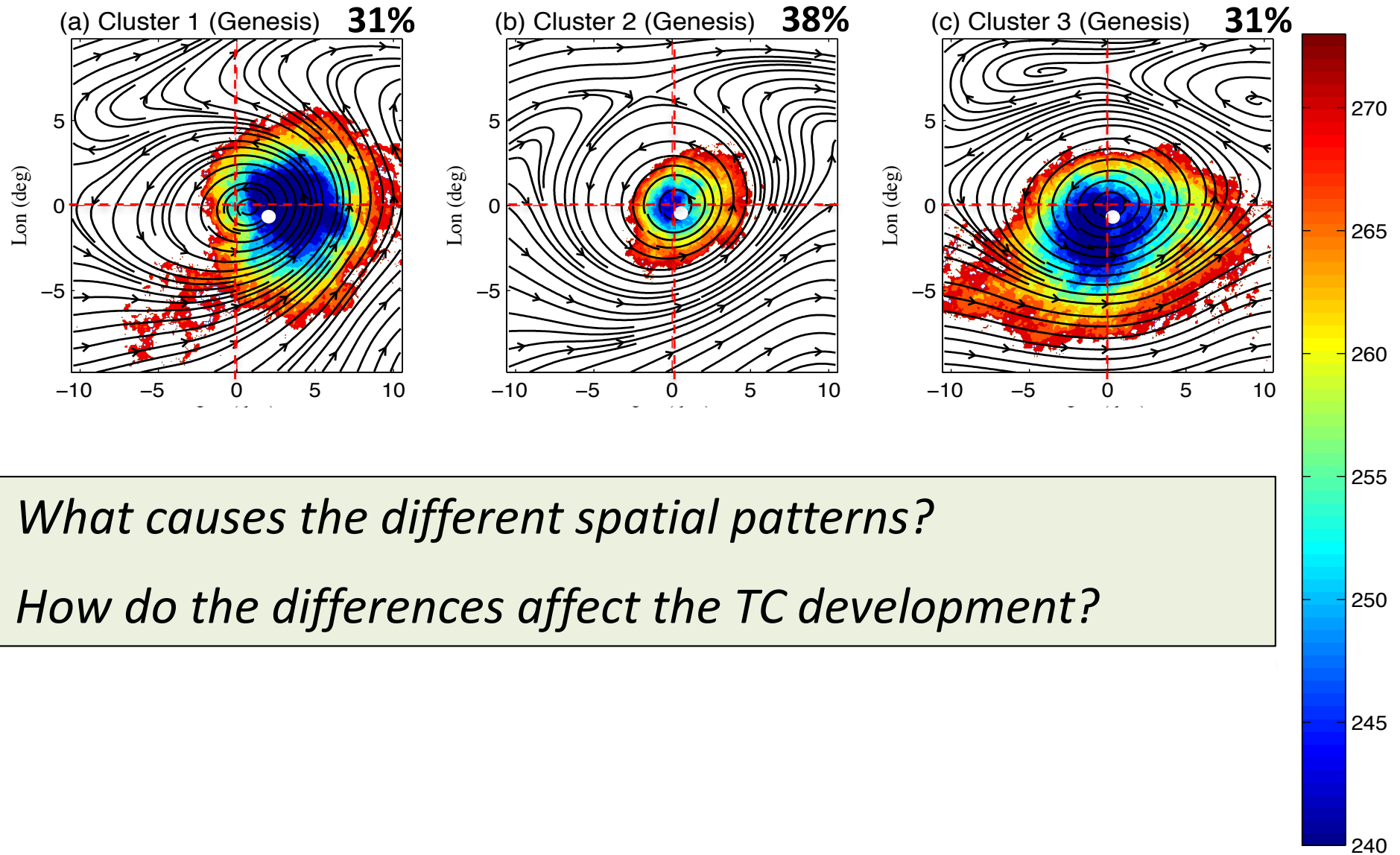
# IR Composites



Composites with respect to the pouch center.

# Three Clusters of Different Spatial Patterns: IR

Genesis Time

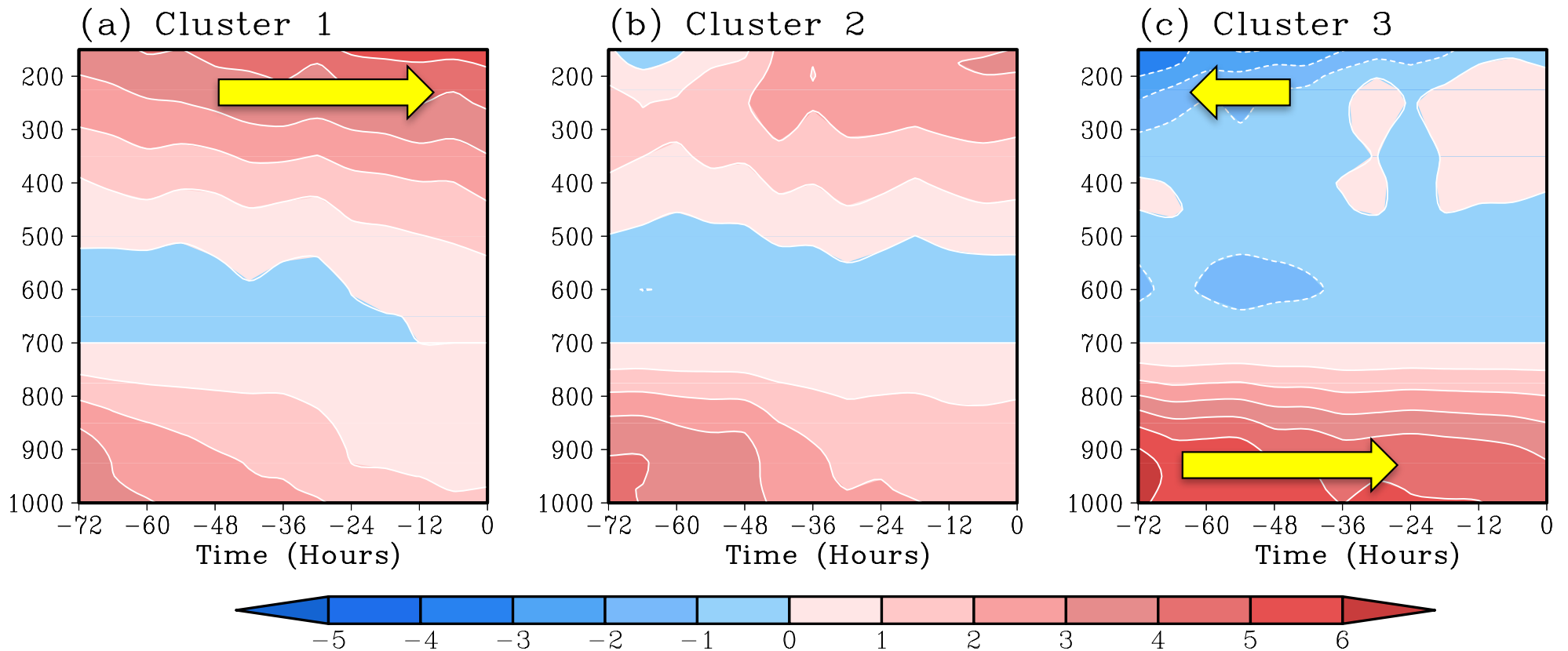


*What causes the different spatial patterns?*

*How do the differences affect the TC development?*

# Three Clusters : Relative Flow (10deg Box)

## $U(p)-U(700\text{ hPa})$

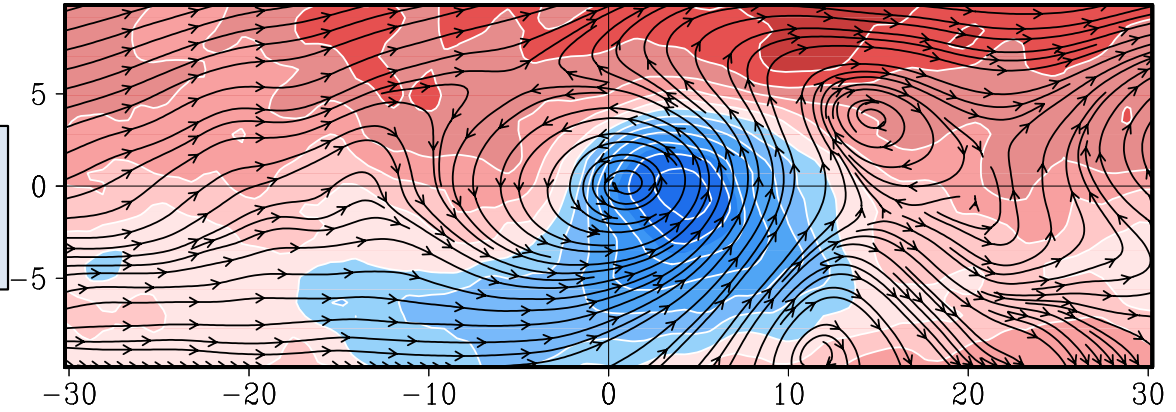


Cluster 1: strong westerly relative flow in the upper troposphere  
Cluster 3: weak easterly relative flow in the upper troposphere;  
strong westerly relative flow in the boundary layer



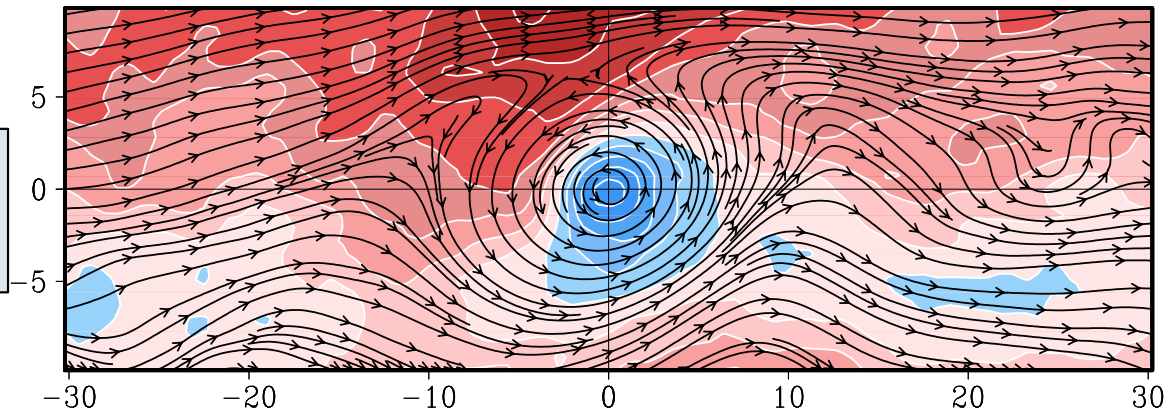
# Three Clusters : Relative Humidity (600 hPa)

(a) Cluster 1



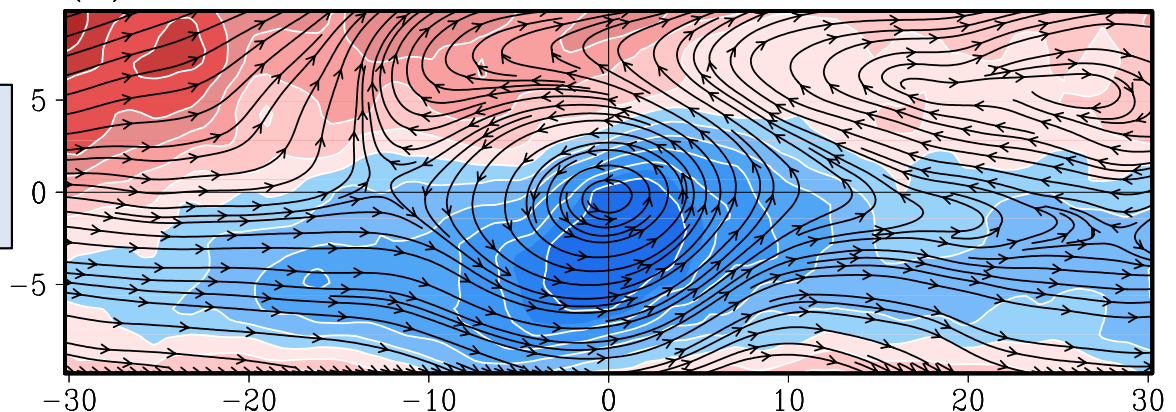
Cluster 1: A moist pouch detached from the ITCZ?

(b) Cluster 2



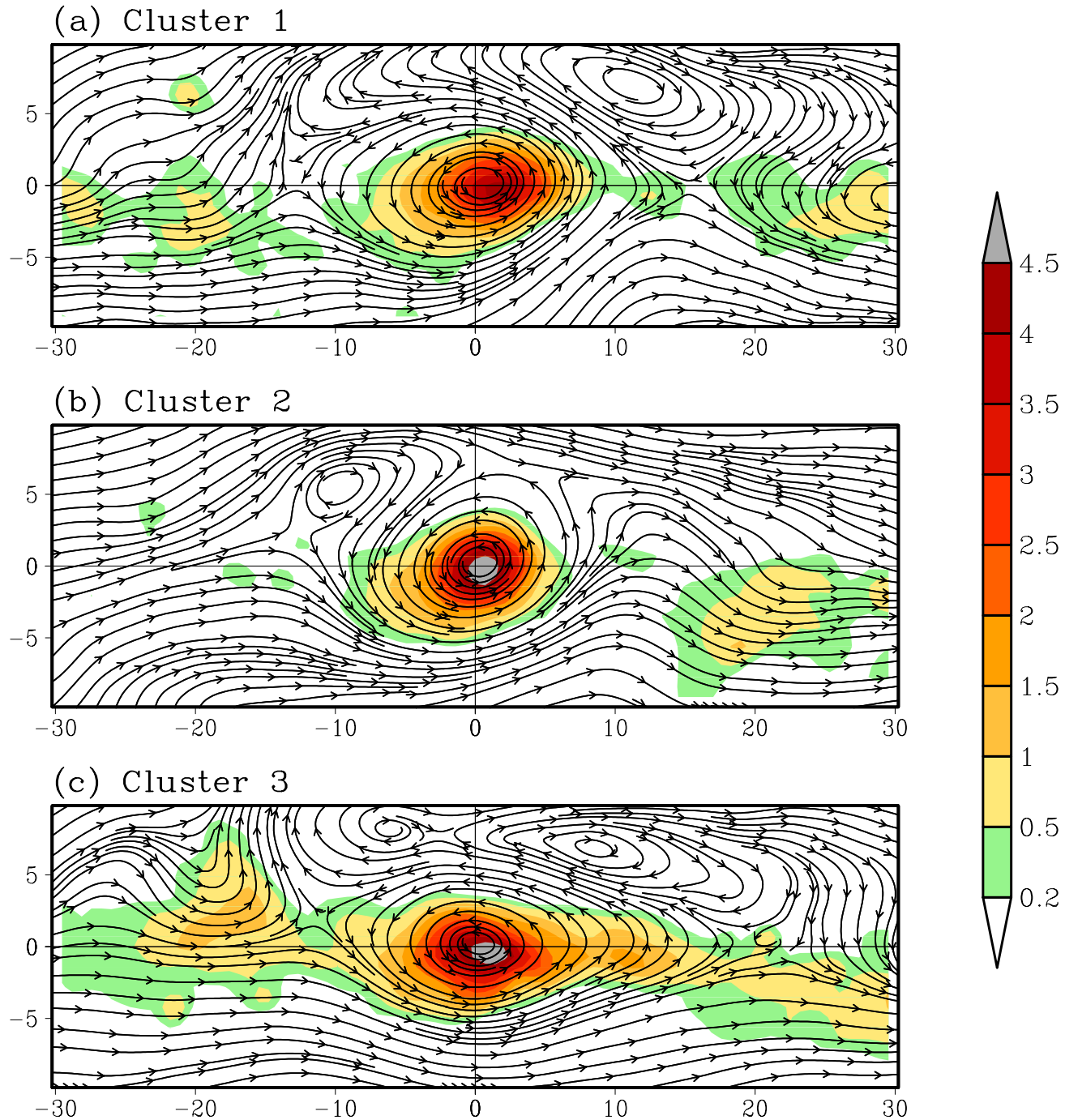
Cluster 2: An isolated moist core surrounded by dry air.

(c) Cluster 3



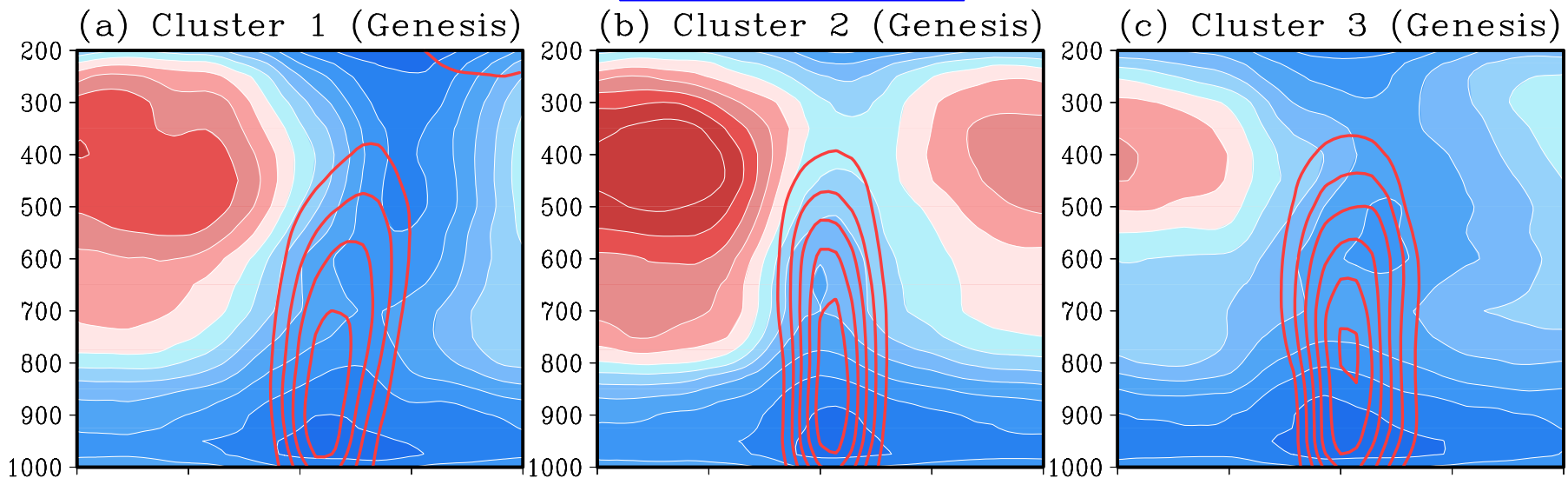
Cluster 3: A large moist pouch embedded in the ITCZ

# Three Clusters : Relative Vorticity (700 hPa)

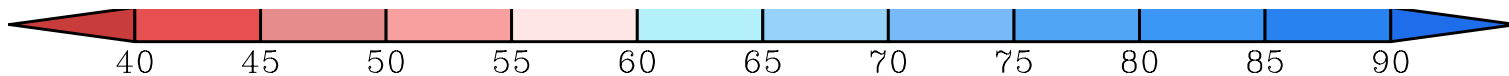


# East-west Vertical Cross Section: OW (contours) and RH

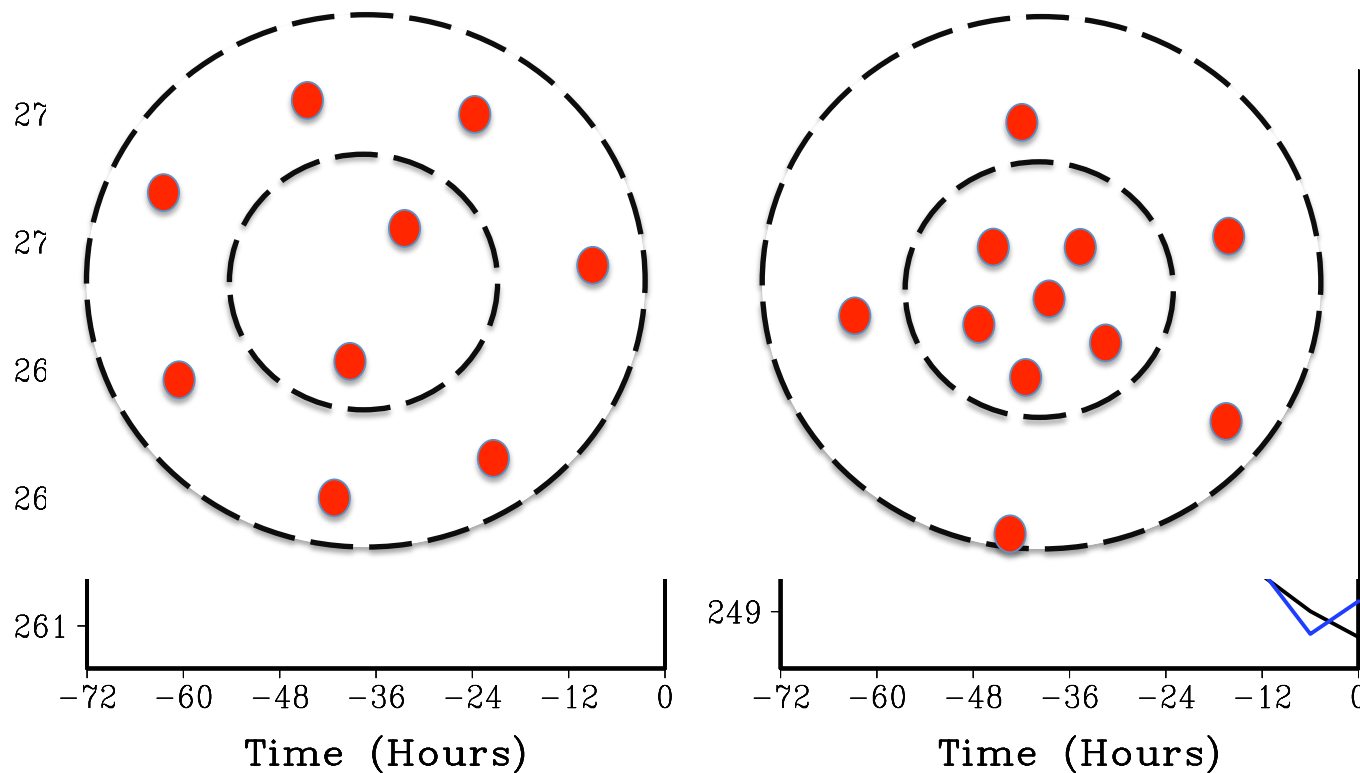
Genesis Time



- A small convective core does not mean a weaker vortex.
- Displacement of convection off the pouch center is associated with a weaker vortex.



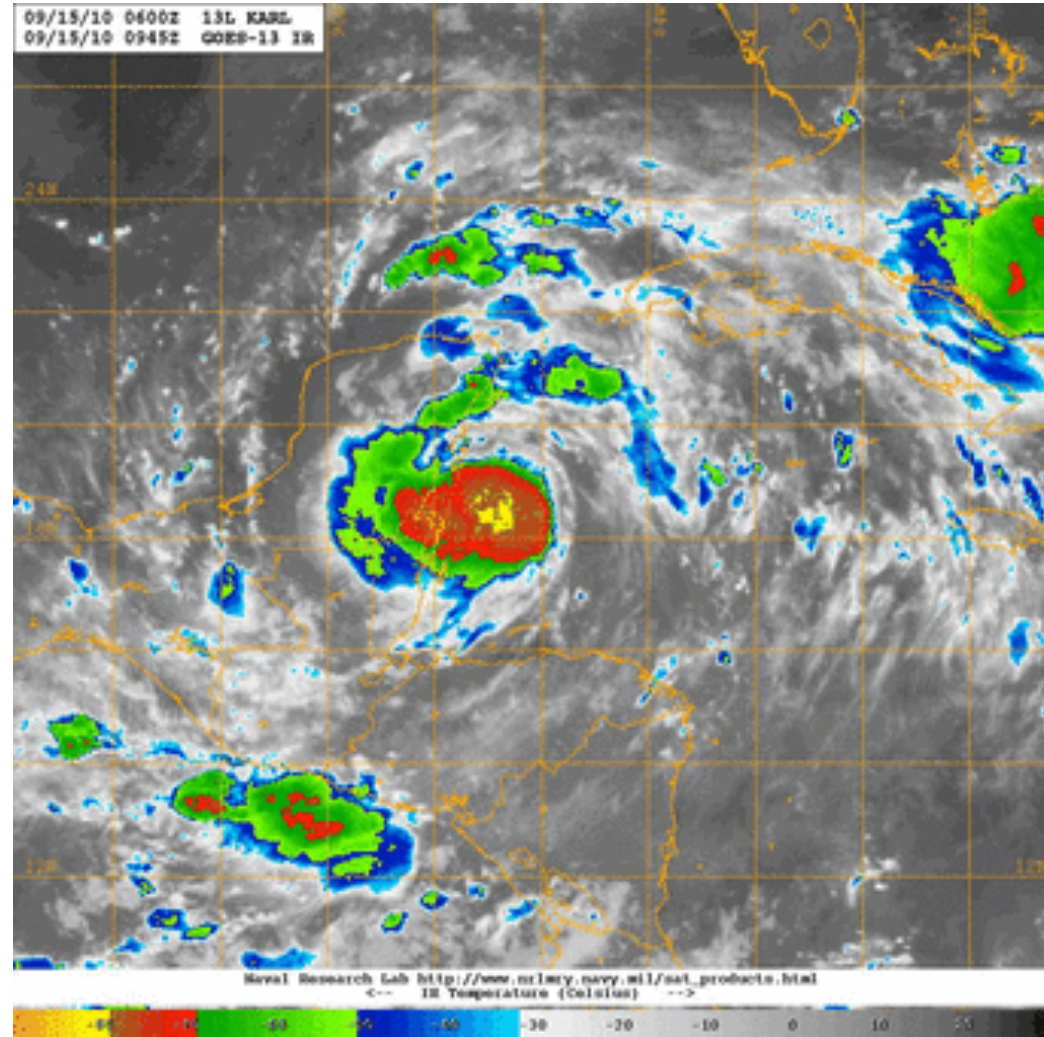
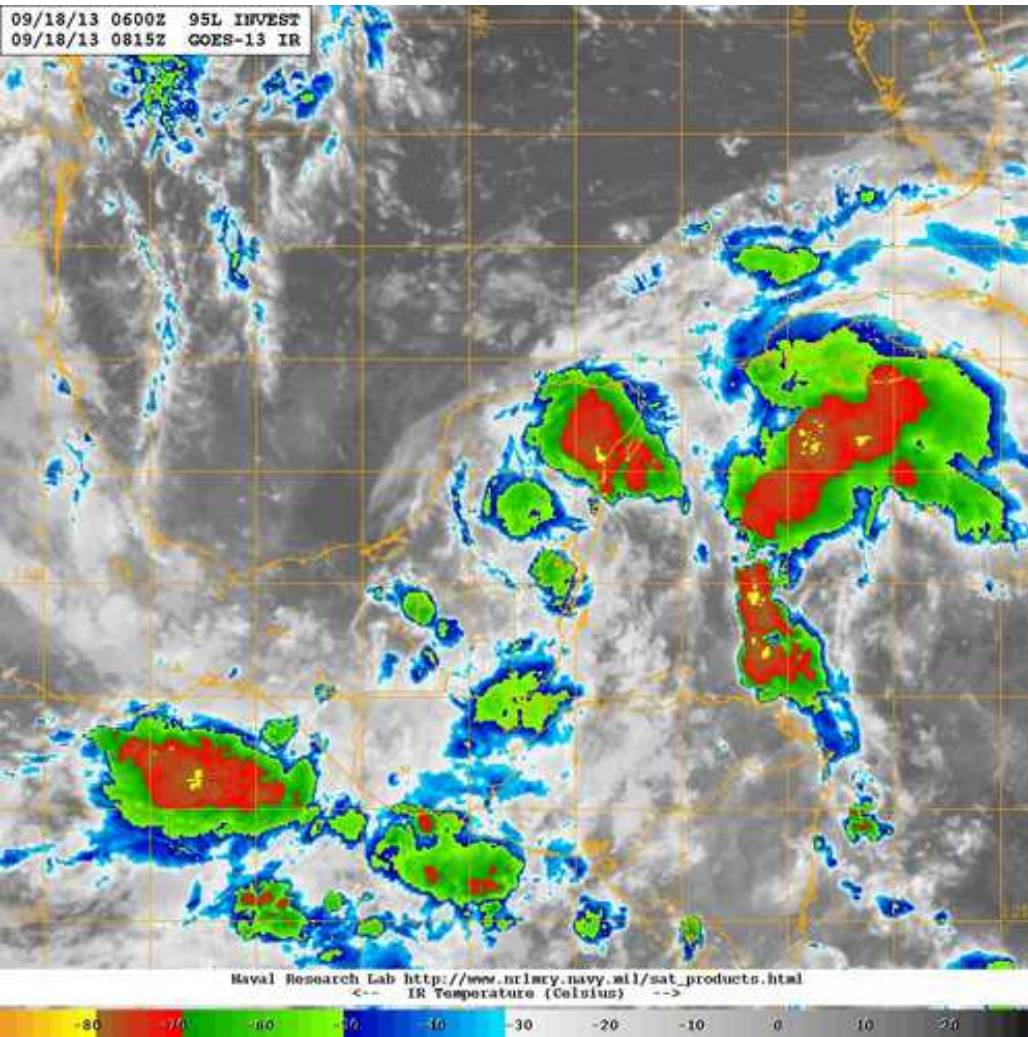
# Time Series of Median IR for Each Cluster



$$\frac{\partial}{\partial r} \left( A \frac{\partial r \psi}{r \partial r} + B \frac{\partial \psi}{\partial z} \right) + \frac{\partial}{\partial z} \left( B \frac{\partial r \psi}{r \partial r} + C \frac{\partial \psi}{\partial z} \right) = \frac{g \partial Q}{\theta \partial r}$$

The transverse circulation is driven by the gradient of diabatic heating, instead of the heating itself! Heating is most effective when the maximum is collocated with the vorticity center.

# IR imagery: which one is a tropical cyclone?



# Three Clusters : TC Size

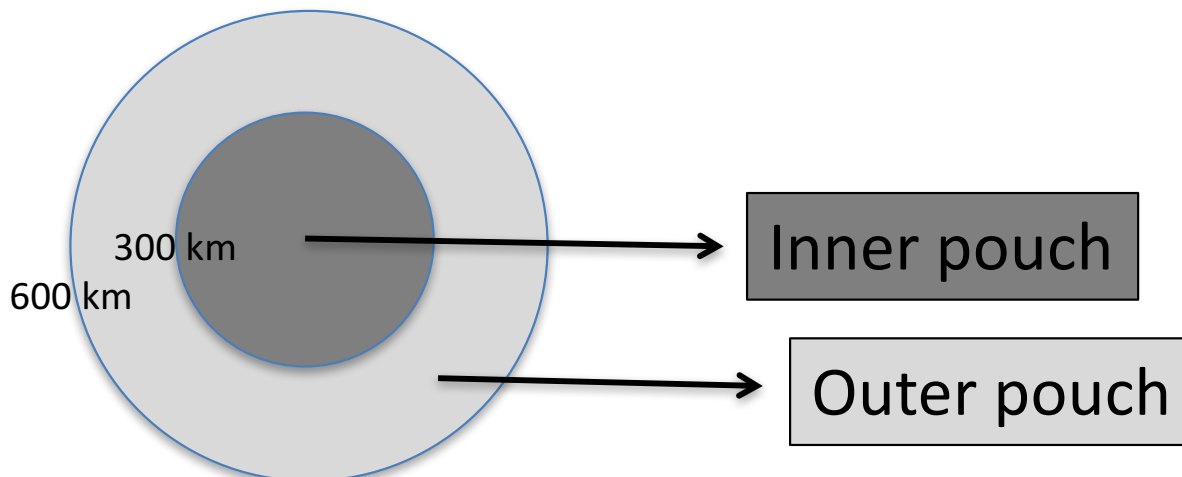
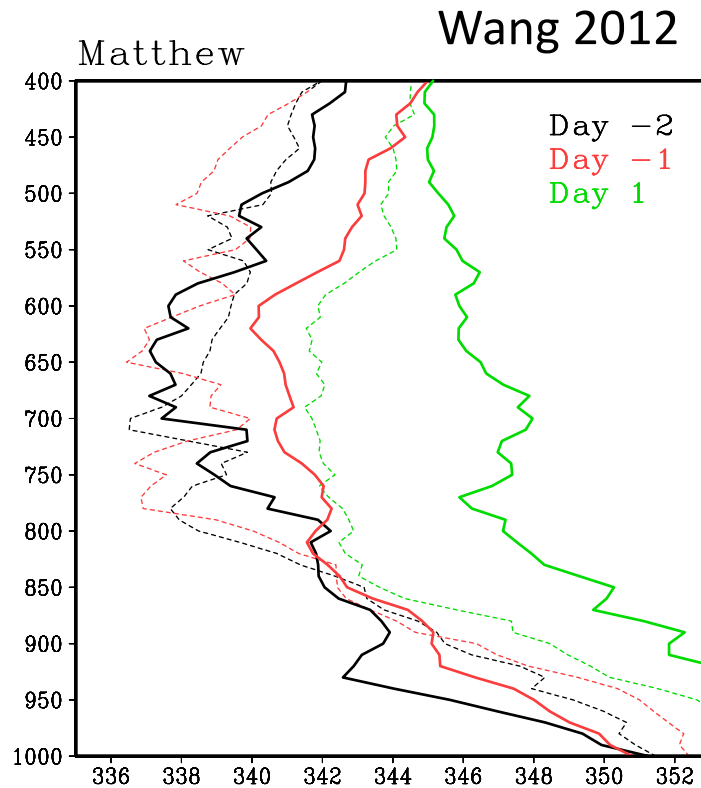
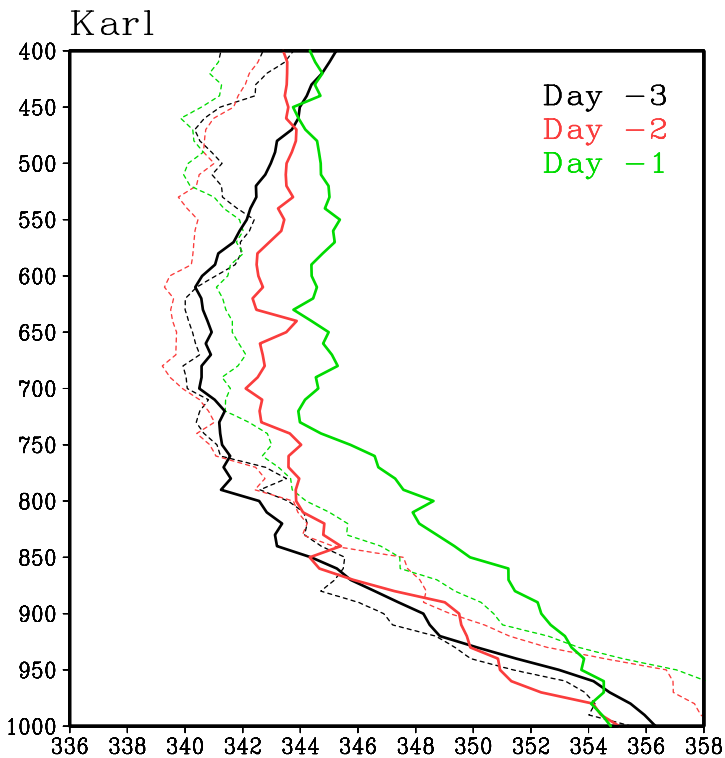
Cluster	1	2	3
Mean AR34	72	59	66
Median AR34	72	53	62

AR34: Radius of 34-knot winds averaged over one day after the declaration of a tropical storm.

Restricted to a latitude band 10-25N.

Q2: What thermodynamic conditions promote convective organization near the pouch center?  
--- SSMI/SSMIS data: simultaneous rain rate and CWV retrievals

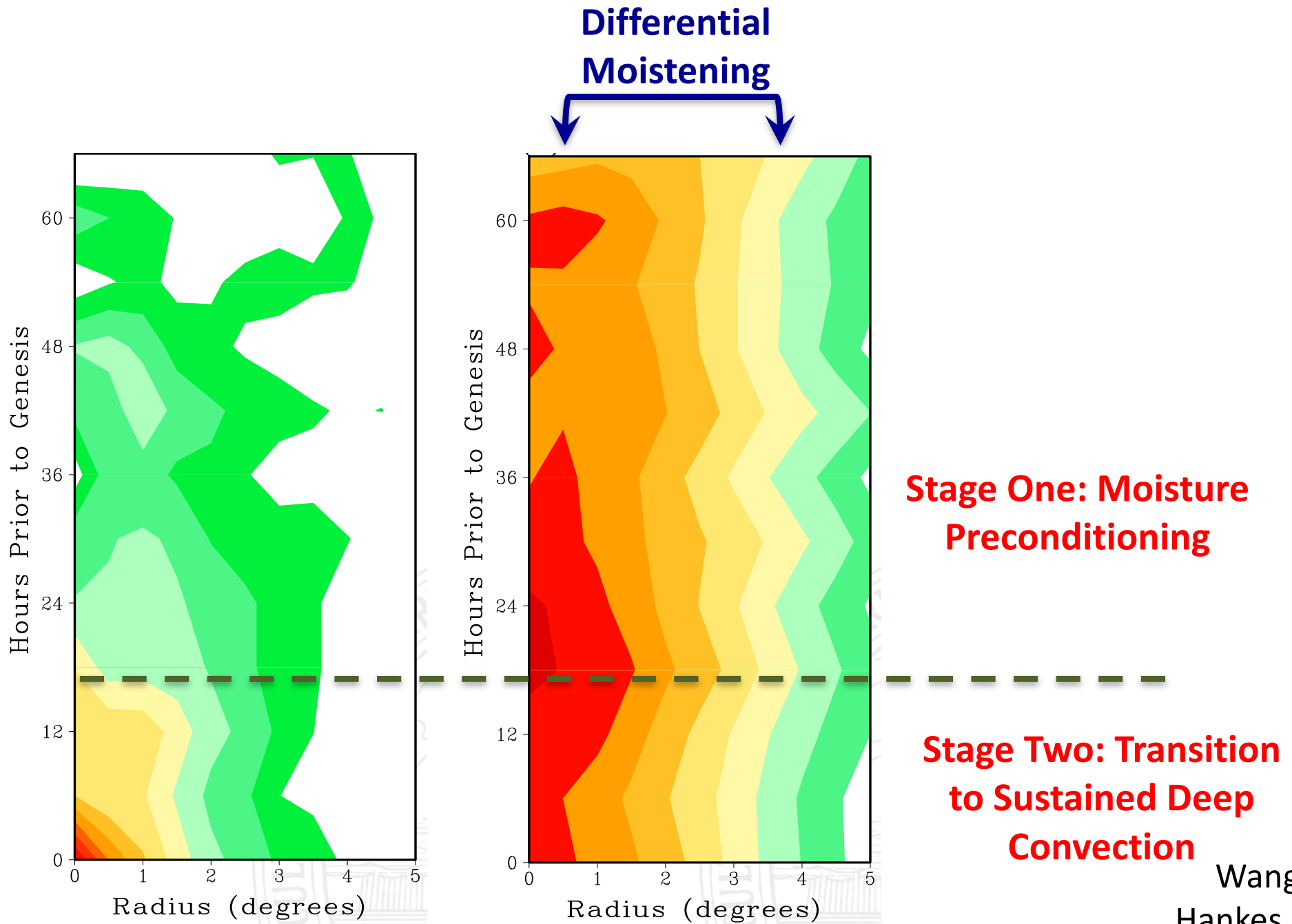
# Equivalent Potential Temperature Profiles: inner vs. outer pouch region



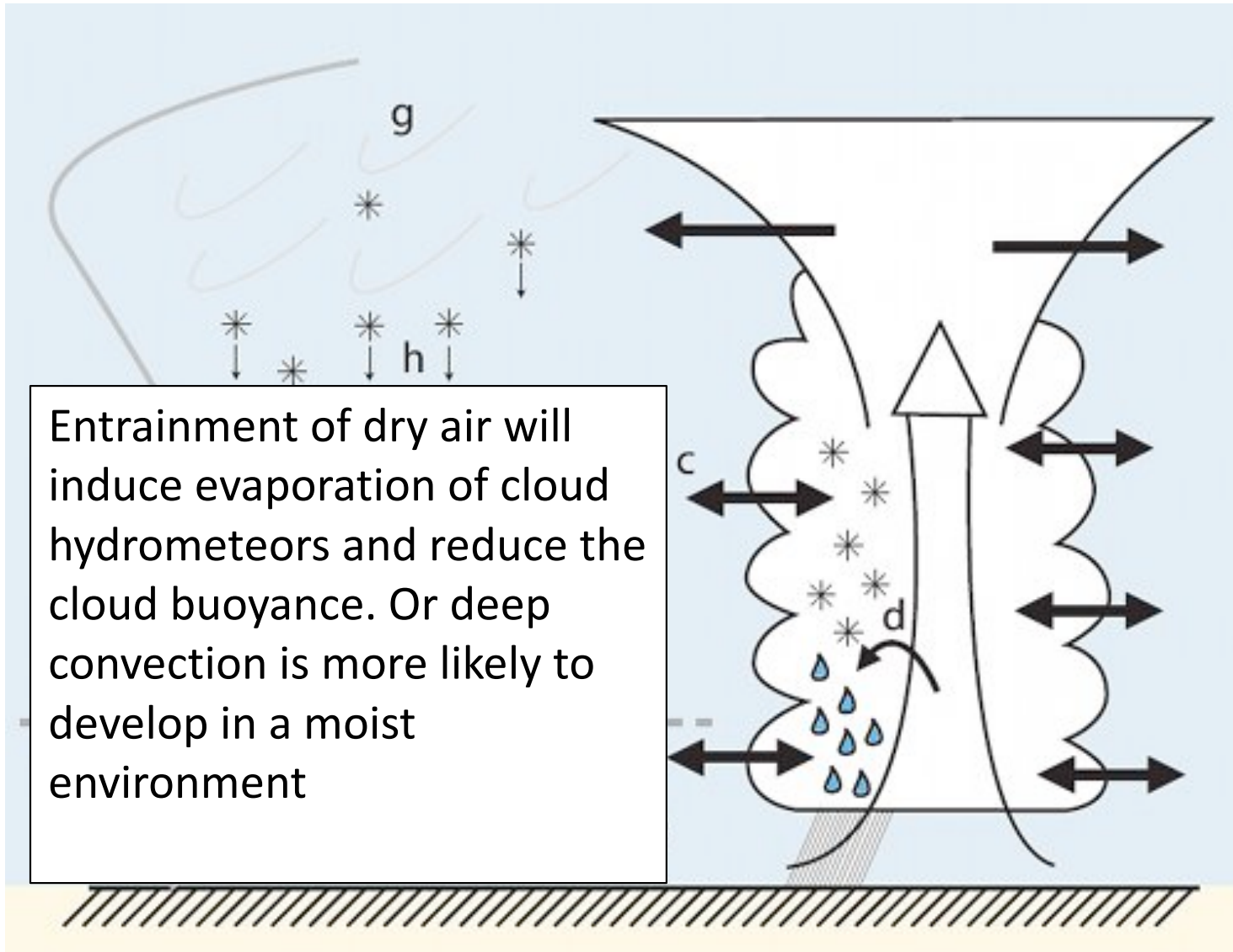
- The midlevel  $\theta_e$  increases significantly near the pouch center one to two days prior to genesis but changes little away from the pouch center.
- This may be an indicator of the impending TC genesis.
- Consistent with Nolan (2007)



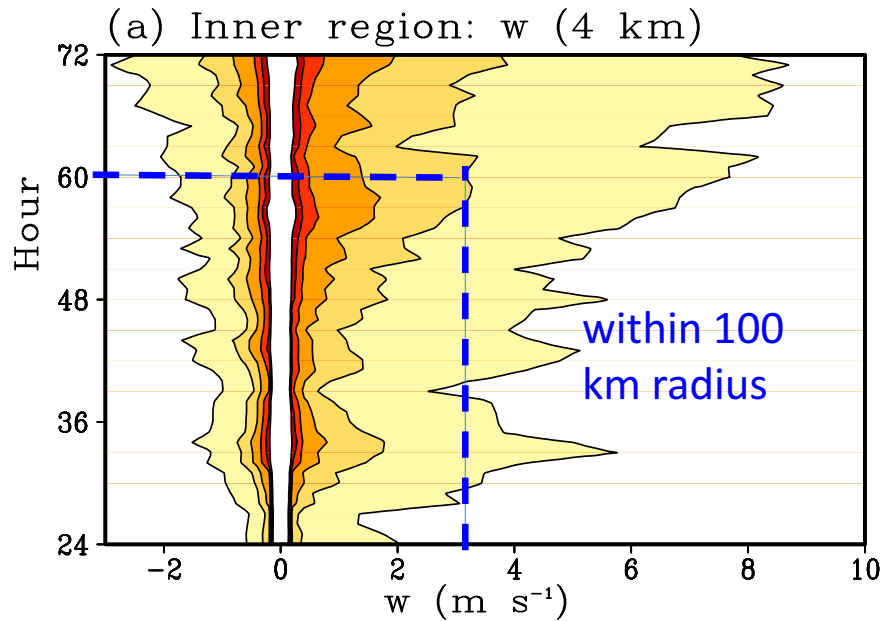
# Time-Radius Plots of Rain Rate and CWV



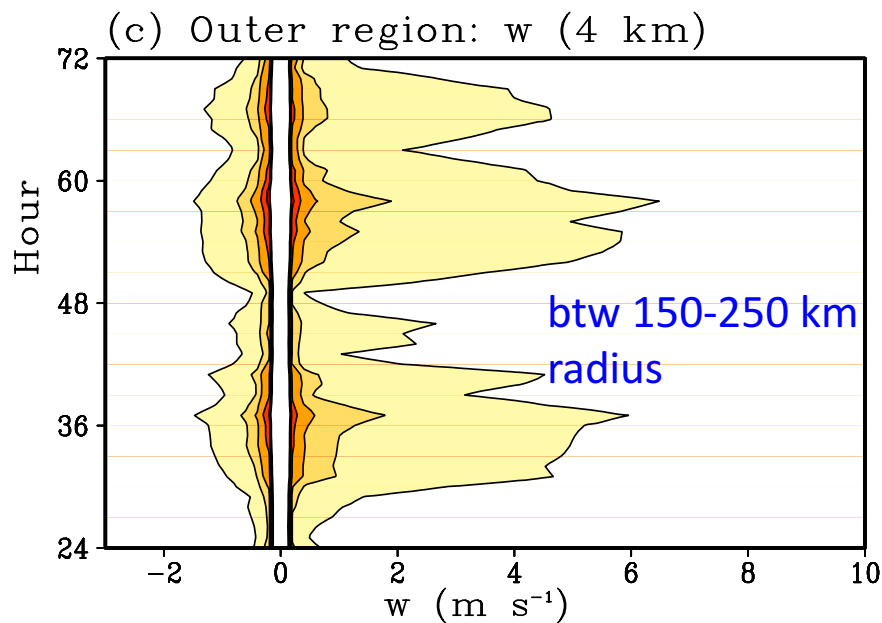
# Dry air and Cloud Buoyancy



# WRF Simulations: Cumulative Distribution of $w$



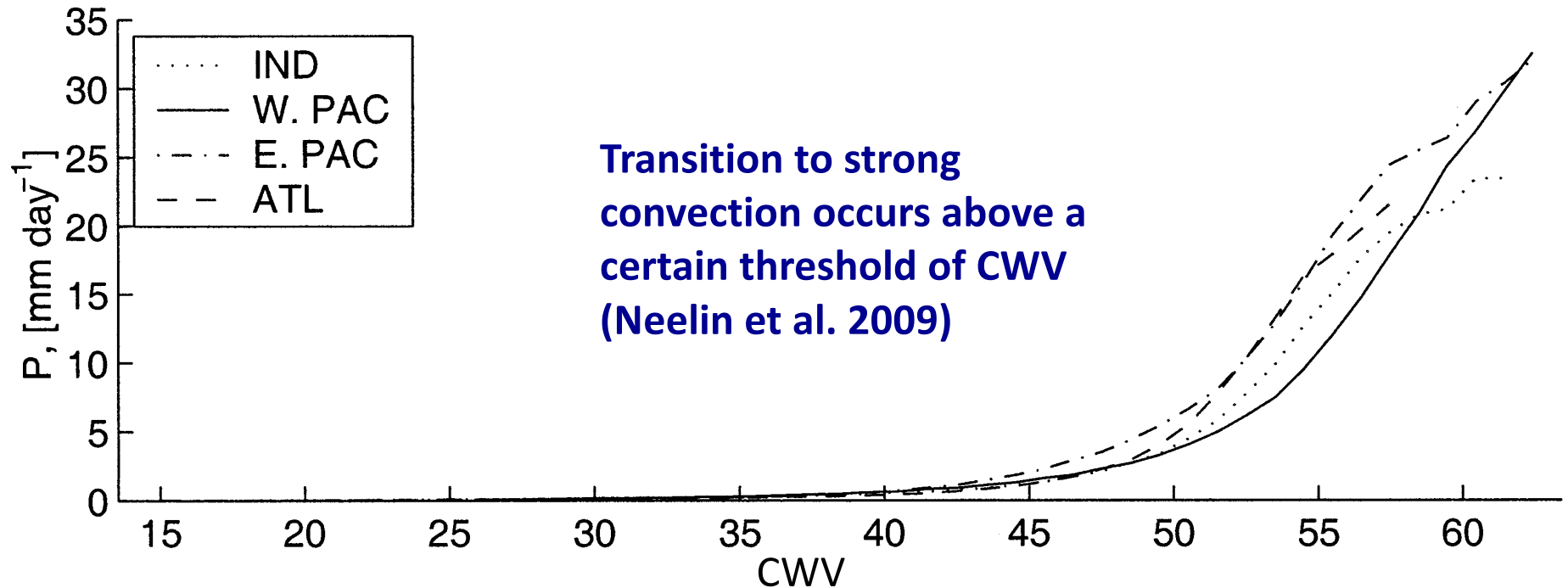
In a nearly saturated column, dry air entrainment is reduced and cloud plumes are more buoyant.



# Precip vs. Column Water Vapor

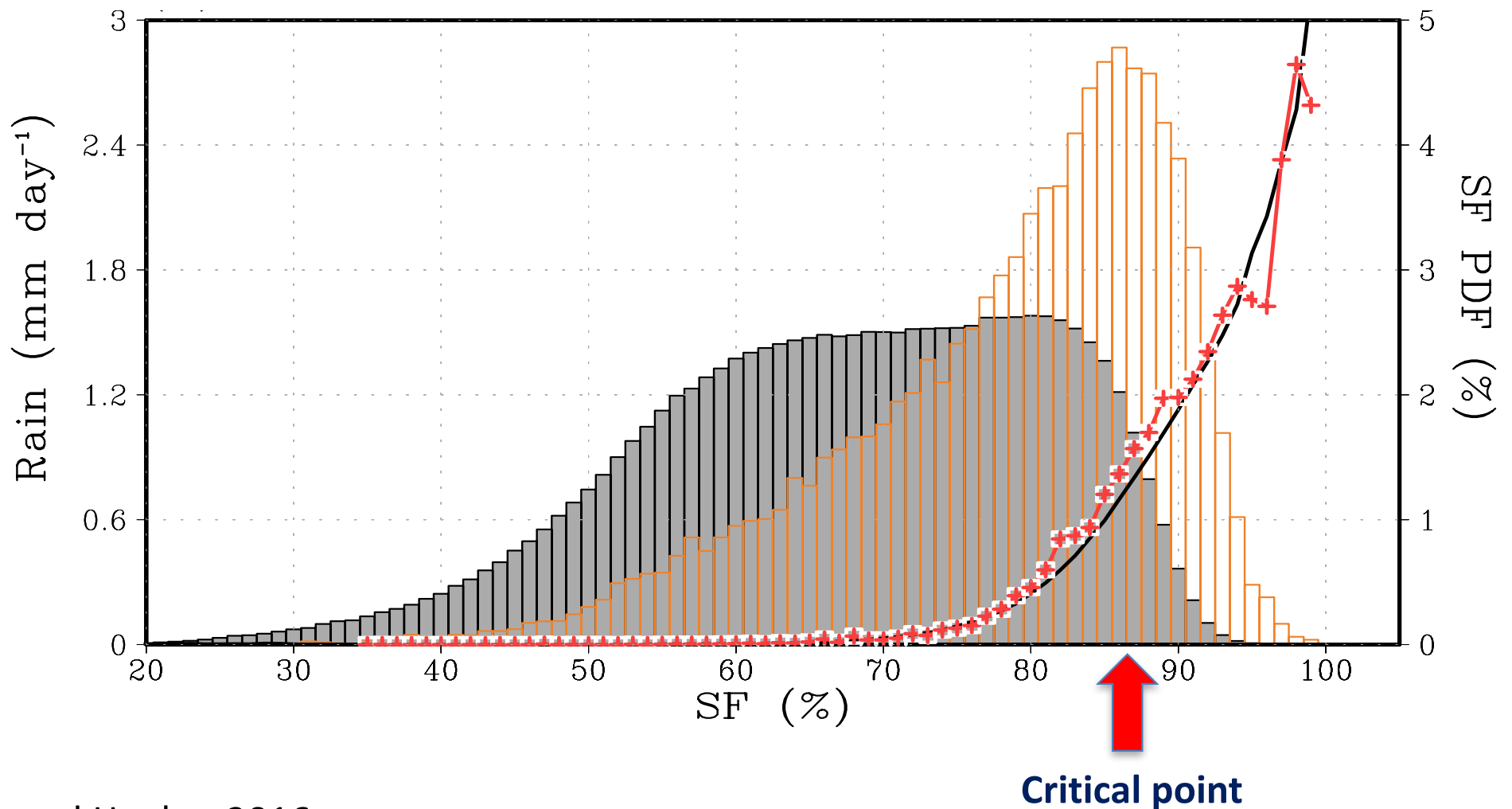
Daily-mean, P vs W

Bretherton et al. 2004



Precip increases exponentially with the column water vapor (Raymond 2000; Bretherton 2004).

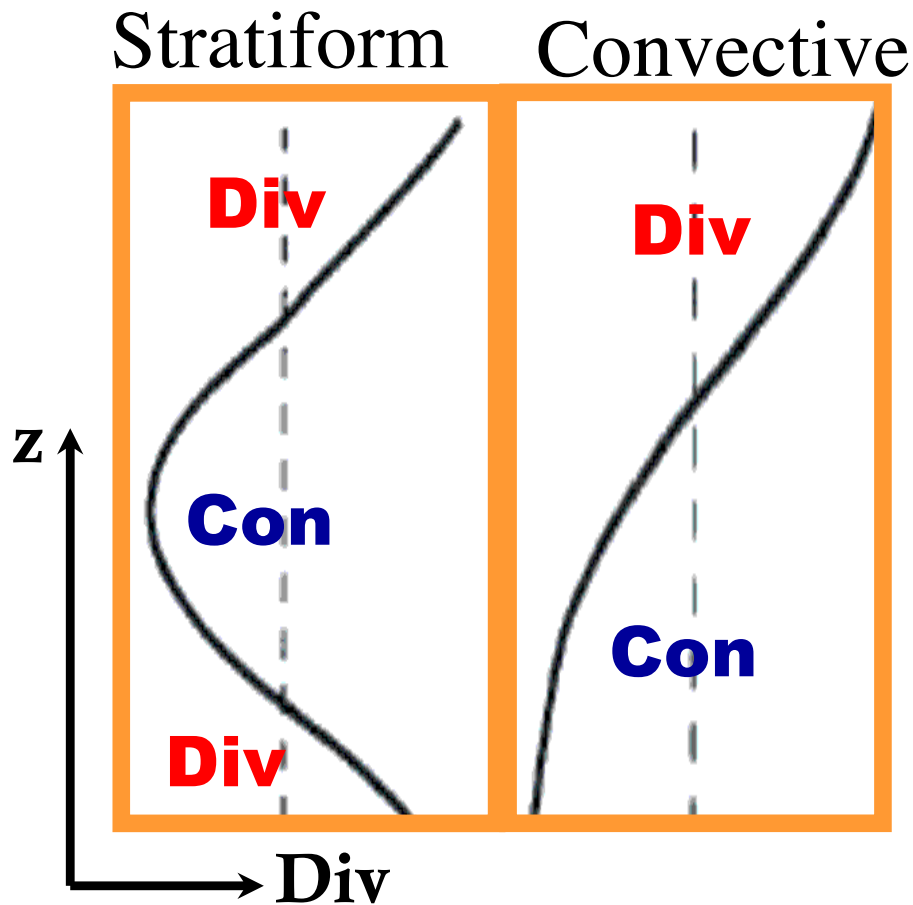
# Column RH-Precipitation relationship



Q3: What are the relative roles of different types of precipitation in TC genesis?

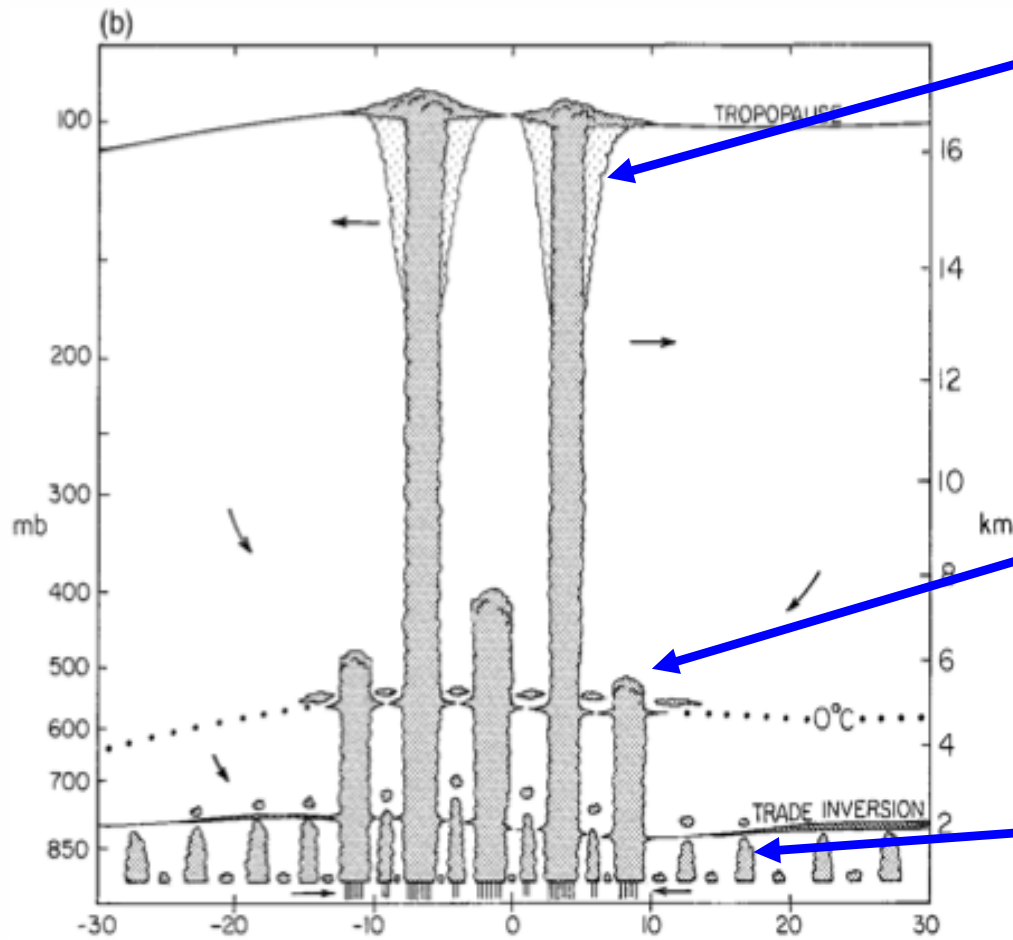
--- TRMM PR 2A25: vertical profile of reflectivity

# Stratiform vs. Convective Precipitation



- Stratiform process: favors the development of a mid-level vortex.
- Convective process: favors the spin-up of the low-level circulation.

# Trimodal Distribution of Convective Clouds



## Partitioning of convection:

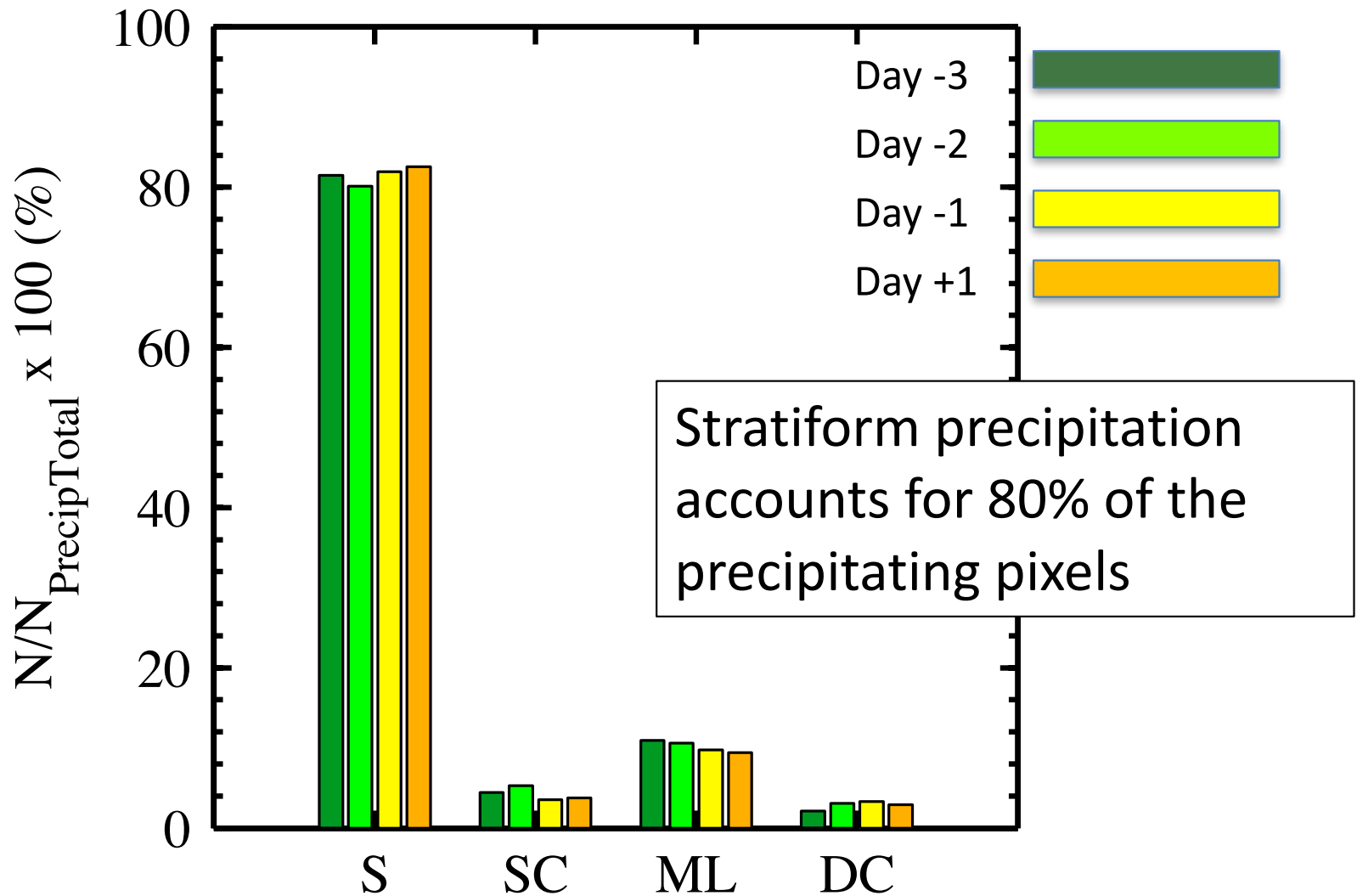
Echo top height derived from 2A25 reflectivity was used to partition the three types of convection.





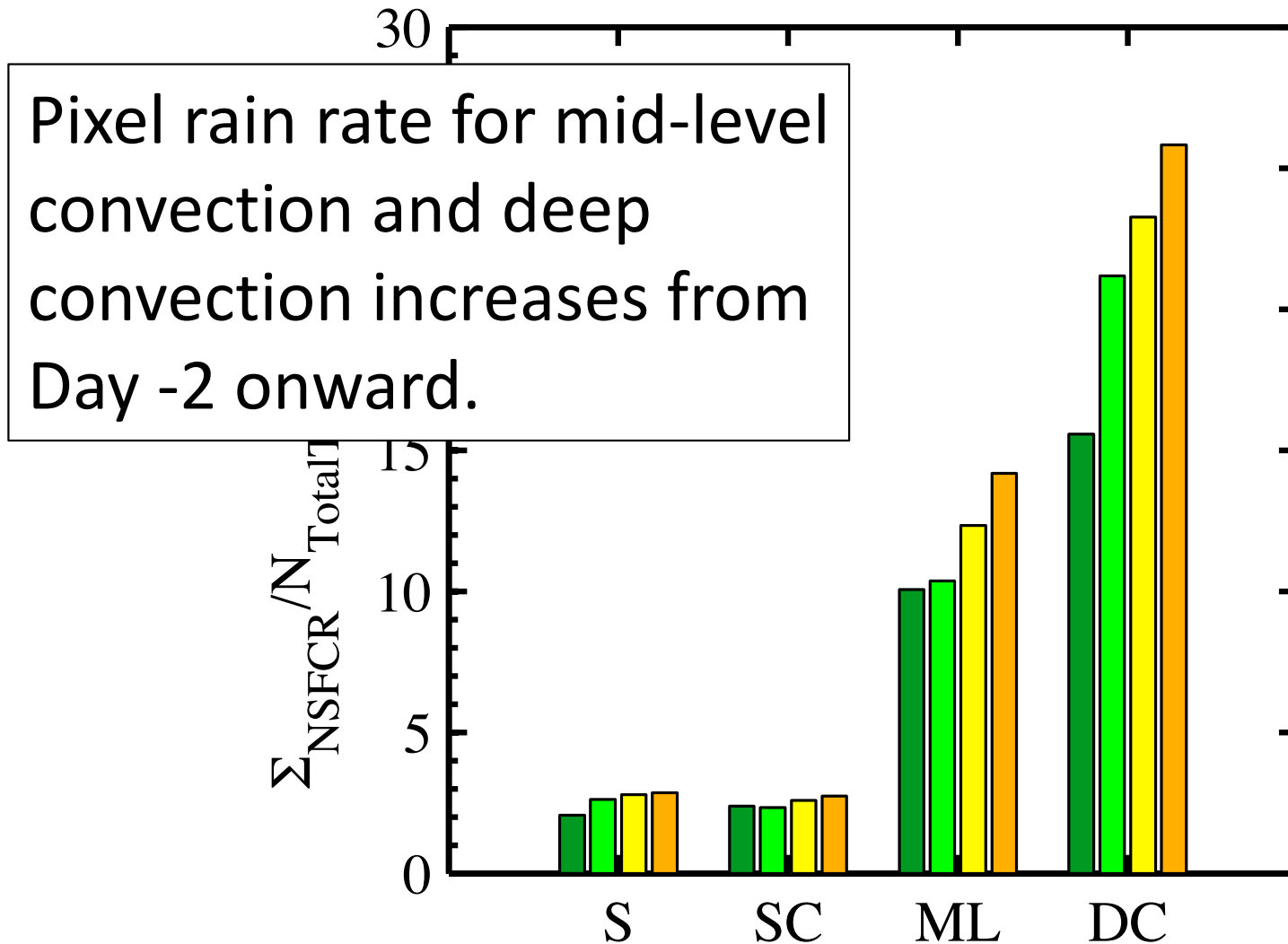
# TRMM PR: Frequency of occurrence of Precipitating Pixels

(c) FO Precipitating Pixels



# TRMM PR: Pixel Rain Rate

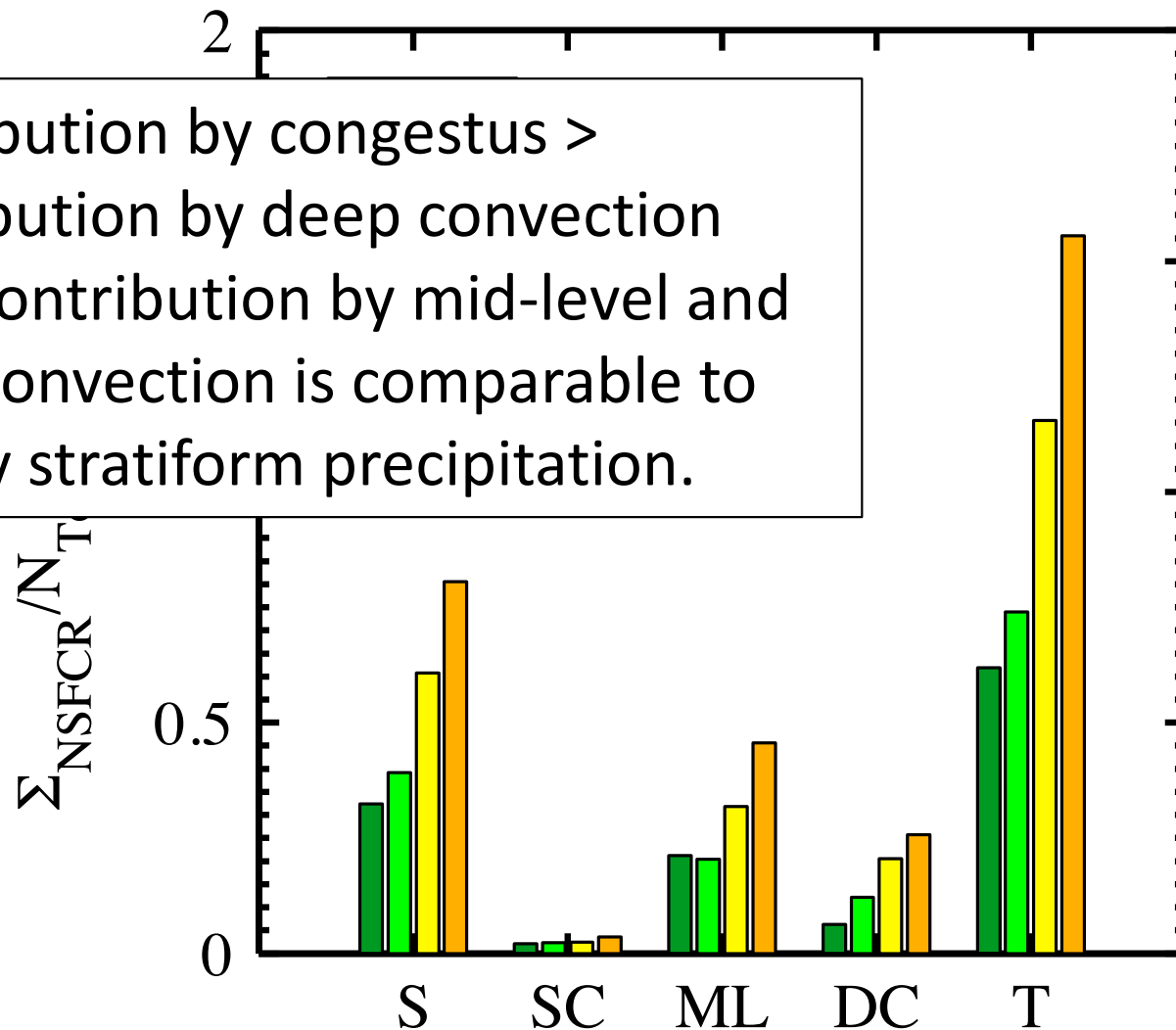
(d) Cond. Mean Rain Rate



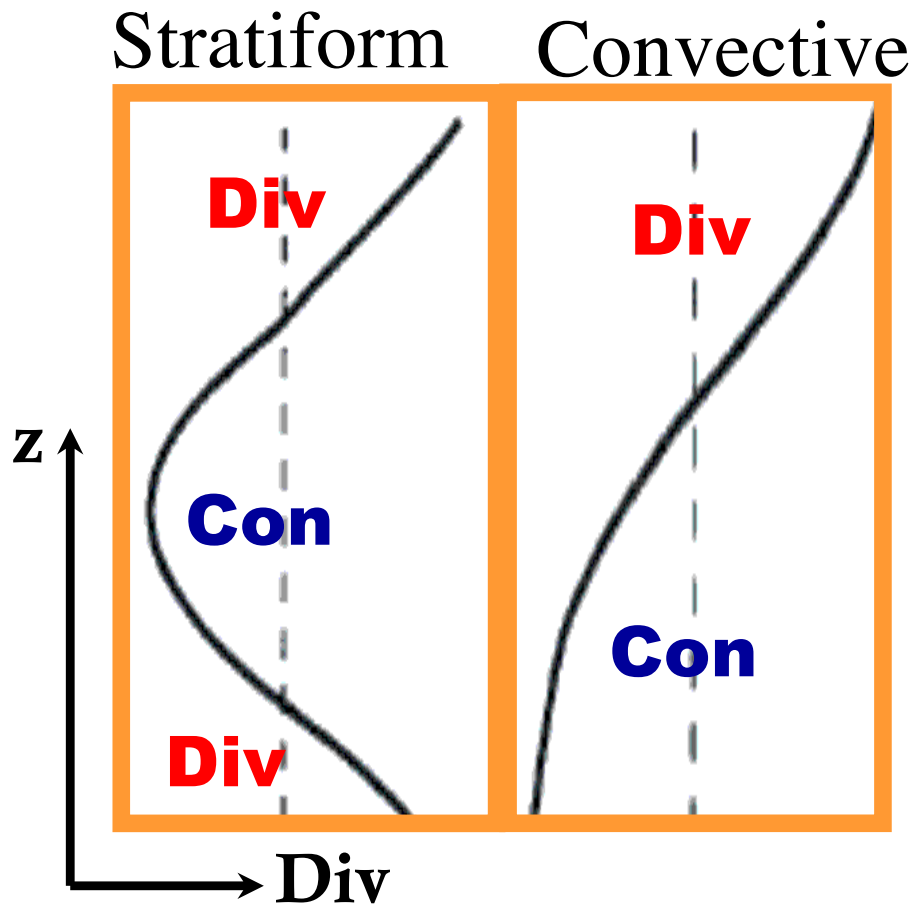
# TRMM PR: Contribution to the Mean Rain Rate

(a) Contrib. Mean Rain Rate

1. Contribution by congestus > contribution by deep convection
2. Total contribution by mid-level and deep convection is comparable to that by stratiform precipitation.

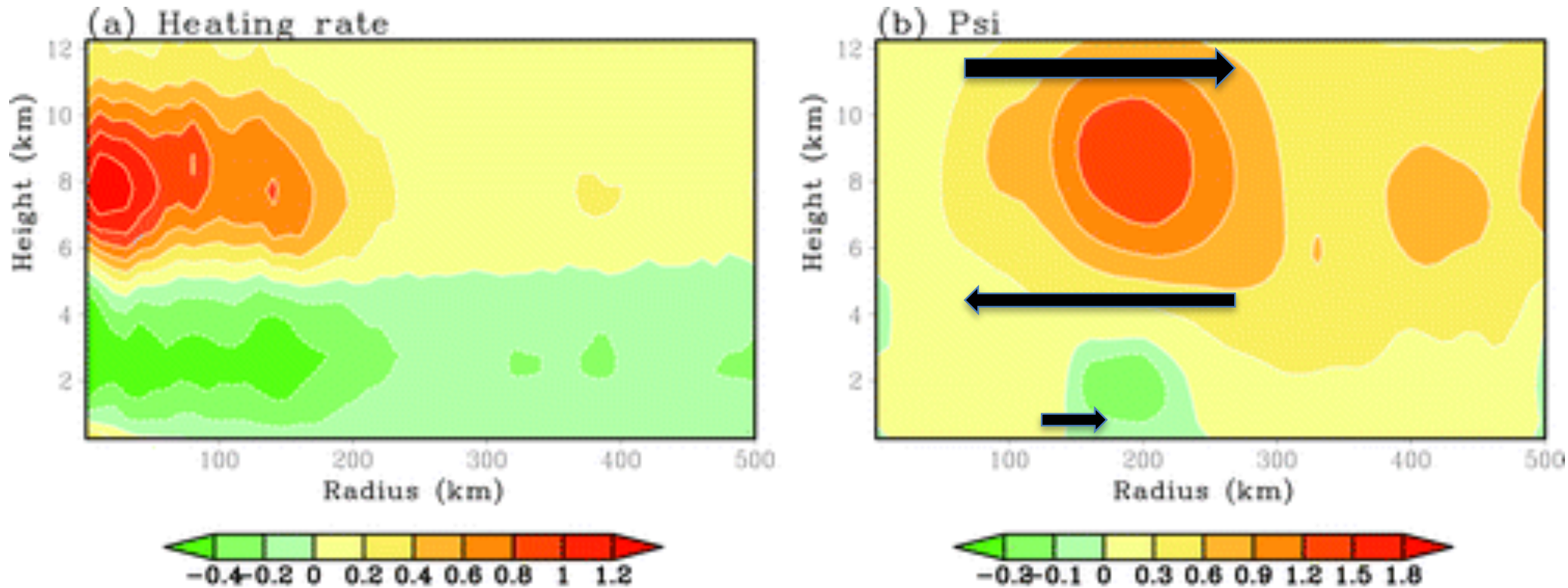


# Stratiform vs. Convective Precipitation



- Stratiform process: favors the development of a mid-level vortex.
- Convective process: favors the spin-up of the low-level circulation.

# Stratiform Precipitation



Stratiform heating contributes to the midlevel spinup without significantly spinning down the low-level circulation.

# Summary

- The key feature of convection for TC genesis is not the intensity or extent of deep convection, but the convective organization near the pouch center.
- Column moistening near the pouch center precedes the transition to sustained deep convection and tropical cyclogenesis.
- Tropical cyclogenesis may be an outcome of the collective contribution by different types of precipitation.